1 2 3 4	Extreme storms cause rapid but short-lived shifts in nearshore subtropical bacterial communities
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15 16	Originality-Significance Statement
17	Extreme storm events, such as tropical cyclones, can negatively affect coastal ecosystems
18	through increased terrestrial run-off, pollution, and physical destruction. Future climate
19	scenarios predict increased frequency and intensity of tropical cyclones, necessitating a better
20	understanding of how ecosystems respond to such events. In this study, we show the short-
21	term dynamics of nearshore bacterial communities during two major tropical cyclones occurring
22	at the beginning and end of the typhoon season in the subtropical western Pacific. Importantly,
23	field observations were coupled with concurrent mesocosm experiments to isolate the effects of
24	terrestrial sediment input from other storm effects, such as wind, waves, and fresh-water influx.
25	Our study reveals that shifts in bacterial communities in both instances were extremely rapid but
26	highly context dependent.
27 28 29	Abstract
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31	Climate change scenarios predict tropical cyclones will increase in both frequency and intensity,
32	which will escalate the amount of terrestrial run-off and mechanical disruption affecting coastal
33	ecosystems. Bacteria are key contributors to ecosystem functioning, but relatively little is known

about how they respond to extreme storm events, particularly in nearshore subtropical regions. 34 35 In this study, we combine field observations and mesocosm experiments to assess bacterial 36 community dynamics and changes in physicochemical properties during early- and late-season 37 tropical cyclones affecting Okinawa, Japan. Storms caused large and fast influxes of freshwater 38 and terrestrial sediment—locally known as red soil pollution—and caused moderate increases of macronutrients, especially SiO₂ and PO_{4³⁻}, with up to 25 and 0.5 μ M, respectively. We 39 40 detected shifts in relative abundances of marine and terrestrially-derived bacteria, including 41 putative coral and human pathogens, during storm events. Soil input alone did not substantially 42 affect marine bacterial communities in mesocosms, indicating that other components of run-off 43 or other storm effects likely exert a larger influence on bacterial communities. The storm effects 44 were short-lived and bacterial communities guickly recovered following both storm events. The 45 early- and late-season storms caused different physicochemical and bacterial community 46 changes, demonstrating the context-dependency of extreme storm responses in a subtropical 47 coastal ecosystem.

48

49 Keywords

Tropical cyclones, typhoons, hurricanes, extreme events, bacterioplankton, coastal, nearshore,
community dynamics, soil pollution, run-off

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53 Introduction

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55 Extreme storm events, such as tropical cyclones (i.e. tropical storms, hurricanes, and typhoons),

56 can have dramatic consequences on coastal ecosystems, due in part to the effects of

57 terrestrially-derived pollution (Hennessy *et al.*, 1997; De Jesus Crespo *et al.*, 2019). In addition

- to influencing salinity and turbidity, flood plumes often include elevated concentrations of
- 59 bacteria (Solo-Gabriele et al., 2000), nutrients (i.e. C, N, P) (Chen et al., 2012, 2018; Gao et al.,

60 2014; Paerl et al., 2018) and other chemicals, such as herbicides or heavy metals (Lewis et al., 61 2012; Mistri et al., 2019), which can act synergistically to negatively affect coastal ecosystems 62 (Wooldridge, 2009; Brodie et al., 2012; Lewis et al., 2012). Especially in tropical and subtropical 63 regions experiencing severe seasonal storms, large volumes of terrestrial run-off entering 64 coastal waters can degrade coastal ecosystems, including coral reefs, through sedimentation or 65 disease (Riegl and Branch, 1995; Philipp and Fabricius, 2003; Voss and Richardson, 2006; 66 Haapkylä et al., 2011; Wilson et al., 2012). Such run-off events can also cause harm more 67 indirectly, through eutrophication, hypoxia (Fabricius, 2005; Altieri et al., 2017) and decreased 68 water quality. As global climate change is expected to enhance the frequency and intensity of 69 extreme storm events (Groisman et al., 2005), it is increasingly important to better understand 70 how such storms impact coastal ecosystem functioning.

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72 The western North Pacific, where there is an average of 27 named storms per year (Wang et 73 al., 2010; Herbeck et al., 2011), is the most active region in the world for tropical cyclones. 74 Landfalling typhoons, which most affect coastal ecosystems, have intensified in the region; the 75 proportion of category 4 and 5 typhoons striking land more than doubled in the last four 76 decades. Current climate models predict continued intensification of landfalling typhoons 77 affecting mainland China, Taiwan, Korea and Japan, indicating these regions will suffer even 78 more storm-caused losses of life, property, and coastal habitat (Mei and Xie, 2016). Okinawa 79 Island—the largest island of the Ryukyu archipelago at the edge of the western North Pacific— 80 is an ideal natural laboratory for studying storm effects on coastal ecosystems (Figure 1). 81 Okinawa's coral reefs have experienced significant declines in recent decades, due in part to 82 increased storm induced run-off and sedimentation (Omori, 2011; Hongo and Yamano, 2013; 83 Harii et al., 2014), which is exacerbated by agricultural practices and large coastal development 84 projects (Omija, 2004; Masucci and Reimer, 2019). The fine-particle, laterite soils with high iron 85 concentrations found in Okinawa and typical to the region are easily suspended and turn coastal

waters a deep, cloudy red color during the frequent tropical cyclones (Supplemental Figure 1)
(Omija, 2004). These events are locally referred to as Red Soil Pollution (Omori, 2011).

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89 While the biological consequences of storm-induced run-off have been investigated for corals 90 and fish species in Okinawa (Hongo and Yamano, 2013; Inoue et al., 2014; Yamazaki et al., 91 2015; O'Connor et al., 2016; Yamano and Watanabe, 2016), less is known about how tropical 92 cyclones and associated run-off affects coastal microbial communities and especially bacteria 93 (Blanco et al., 2008). Microbial communities contribute to marine ecosystems through primary 94 production and by recycling dissolved organic carbon and nutrients through the microbial loop 95 (Azam et al., 1983), but can also draw down dissolved oxygen (Anderson and Taylor, 2001) or 96 cause opportunistic infections in marine organisms (Shinn et al., 2000; Sutherland et al., 2011; 97 Sheridan et al., 2014; Peters, 2015). Therefore, changes in microbial community compositions 98 in response to storms could precipitate large-scale ecosystem effects. Microbial responses can 99 occur extremely quickly; Gammaproteobacteria, Flavobacteria and many Alphaproteobacteria 100 can increase in abundance within hours when exposed to high nutrient concentrations, whereas 101 the entire microbial community-including archaea, protists, and viruses-can turn over on the 102 scale of less than one day to about a week (Fuhrman et al., 2015). Rapid microbial response 103 times to changing environmental conditions make microbes valuable early-warning bioindicators 104 (Glasl et al., 2017; Pearman et al., 2018), but also hinders their study. Sampling at the scale of 105 microbial response times during tropical cyclones is often dangerous and is further complicated 106 by the poor predictability of storm tracks and intensities (Zhou et al., 2012).

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In this study, we characterize nearshore bacterial community dynamics in response to tropical
cyclones affecting Okinawa Island and isolate the effects of sediment input through controlled
mesocosm experiments. The study included tropical storm Gaemi at the start of the 2018
Okinawa typhoon season (June 16) and successive category 5 super typhoons, Trami and

112 Kong-Rey, on September 30 and October 5, towards the end of the 2018 season. We evaluated 113 physicochemical properties and bacterial community compositions in seawater samples 114 collected before, during/between, and after storms in June and October and in samples taken 115 from mesocosms with and without red soil amendment. The specific aims for this study were to: 116 i) assess how bacterial community composition and physicochemical parameters respond in 117 time to tropical cyclones and sediment input, ii) evaluate the speed of the responses and 118 recovery, and iii) identify potential ecosystem consequences due to extreme storms and 119 sediment input.

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121 Results

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123 Physicochemical responses to extreme storm events and red soil input

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125 Two major storm events affecting Okinawa Island were monitored for this study. The first event 126 included tropical storm Gaemi, which made landfall on June 16 and was the first tropical cyclone 127 of the 2018 typhoon season. Gaemi brought 133.72 mm day⁻¹ of precipitation and maximum 128 wind speeds of 12.1 m s⁻¹ to Okinawa (Figure 2). Substantial precipitation and wind were also recorded during the two days leading up to Gaemi's landfall (69.77 and 83.99 mm day⁻¹ 129 130 precipitation and 10.2 and 12.1 m s⁻¹ wind intensity on June 14 and 15, respectively), but no 131 additional rain was recorded until July (Figure 2). The second event included two category 5 132 super-typhoons, Trami and Kong-Rey, which impacted Okinawa in rapid succession towards the 133 end of the 2018 season (Sept. 29 and Oct. 5). Trami and Kong-Rey caused accumulated rainfall 134 of 239.24 and 87.89 mm day⁻¹ and maximum wind speeds of 26.9 and 12.7 m s⁻¹, respectively. 135 The June and October events represented the two largest rainfalls during the 2018 typhoon 136 season, and typhoon Trami recorded the most extreme sustained and gusting wind speeds in 137 2018 (Figure 2). The extreme winds accompanying Trami and Kong-Rey also caused large

138 waves; Trami and Kong-Rey both brought waves with heights greater than nine meters (Japan

139 Meteorological Agency, <u>https://www.data.jma.go.jp/gmd/kaiyou/shindan/index_wave.html</u>).

140 Temperature (°C), salinity (‰) and turbidity (NTU) were measured in situ at the four nearshore 141 field sites (A1–4, Figure 1) at the same time that water samples were collected. There was a 142 significant decrease in salinity (~15‰) and concurrent increase of turbidity (~10-fold) during the 143 storm on June 16 compared to before the storm on June 13 and afterwards on June 19 (Figure 144 3). Up to a 5‰ decrease in salinity and 3 NTU increase in turbidity was measured between the 145 two storms in October (Oct 1 and 3). These changes were smaller in magnitude than were 146 recorded for the June storm and were not statistically significant, despite the storms in October 147 delivering much more precipitation than the June storm (Figure 2). However, the wind 148 accompanying the October storms was also much more intense (Figure 2), which presumably 149 mixed the water column more thoroughly and prevented freshwater lenses from forming, thus 150 contributing to diminished changes in salinity and turbidity being observed in October compared 151 to June.

152 Concentrations of dissolved nutrients—including Nitrate (NO₃), Nitrite (NO₂), Ammonium 153 (NH₄⁺), Phosphate (PO₄³⁻), Silica (SiO₂), and dissolved Iron (dFe)—were measured in seawater 154 samples collected in June and October (Figure 3) and throughout the second mesocosm 155 experiment (Supplemental Figure 6). Overall, dissolved nutrient concentrations were higher in 156 June than October, with the exception of dFe and SiO_2 , which both had similar values in the two 157 sampling months (Figure 3). A nonparametric Kruskall-Wallis test was applied to detect 158 significant storm-induced differences in field nutrient concentrations (Supplemental Table 3). 159 Results showed significant increases (p < 0.05) in NO₂⁻ concentrations during and following storm events in June and October, whereas SiO₂ and PO₄³⁻ were only significantly elevated 160 161 during and after the storm in June (Figure 3).

162 Red soil addition in the October mesocosm experiment caused a significant increase in SiO_2 163 concentration (Supplemental Figure 6, Supplemental Table 4). Additionally, red soil addition 164 caused PO_4^{3-} concentration to increase above the detection limit 4 hours following soil addition, 165 whereas the concentration of PO₄³⁻ stayed below the detection limit in control mesocosms after 166 the initial measurement (Supplemental Figure 6). Two-way ANOVA results (Supplemental Table 4) indicate that time had a greater effect on nutrient concentration (p < 0.05 for NO₂, SiO₂ and 167 168 dFe) than red soil treatment (p < 0.05 for SiO₂ and dFe). The treatment by time interaction was 169 only significant in the case of SiO₂ (p < 0.05), for which higher concentrations were found in soil-170 treated mesocosms.

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172 Bacterial community responses to storm events and red soil input

173 Metabarcoding with the bacterial 16S ribosomal RNA gene was performed to evaluate 174 shifts in bacterial community composition associated with extreme storms in the field and with 175 sediment input in mesocosms. There was a clear shift in the relative abundances of bacterial 176 phyla in field samples collected during the June storm compared to before or afterwards. 177 Several phyla that were also present in soil samples became more abundant in the water 178 samples collected during the storm-including Acidobacteria, Actinobacteria, Chloroflexi, 179 Firmicutes, Rokubacteria and Verrumicrobia—but phyla not detected in soil samples also 180 became more abundant during the storm, most notably Epsilonbacteraeota (Figure 4). Principal 181 component analysis (PCA) of Aitchison distances between bacterial community compositions in 182 water samples also clearly illustrated a shift in community composition during the June storm: 183 samples collected during the storm separated along the primary axis (PC1) from the samples 184 collected both before and afterwards (Figure 5). The shift in community composition during the 185 storm was accompanied by an increase in ASV richness (Amplicon Sequence Variants; Figure

6); estimated ASV richness was significantly higher in samples collected during the June stormcompared to samples collected before and afterwards.

188 Soil addition to June mesocosms also influenced bacterial community composition and 189 richness, although to a lesser extent than the storm-influenced nearshore communities. 190 Bacterial phyla dominant in soil samples were detectable in mesocosm samples taken one hour 191 after soil addition, but their relative abundance diminished in samples taken 24 hours later 192 (Figure 4). Likewise, ASV richness increased in samples taken from mesocosms following soil 193 addition and decreased over time (Supplemental Figure 8). Overall, mesocosm incubation time 194 had a larger effect on bacterial community composition than soil addition (Figure 4, 195 Supplemental Figure 9), despite mesocosm conditions (temperature, salinity, dissolved oxygen) 196 remaining similar to ambient conditions throughout the experiment (Supplemental Figures 4–5). 197 In October, the community composition also shifted between the two storms (Trami and 198 Kong-Rey) relative to before and after the overall event window. Additionally, some phyla 199 present in soil samples, particularly Firmicutes, increased in relative abundance within the event 200 window, but similar to in June, phyla undetected in soil samples also increased in relative 201 abundance (e.g. Epsilonbacteraeota and Fusobacteria). However, unlike in June, 202 Actinobacteria, Planctomycetes, and Verrucomicrobia were already detectable in water samples 203 collected before Typhoon Trami affected Okinawa (Figure 4). PCA further demonstrated a shift 204 in community composition in samples collected between the October storms compared to 205 before and afterward; samples collected between the two storms separated on the primary axis 206 from samples collected before and after the storms (Figure 5). Interestingly, the estimated ASV 207 richness was lower (by about half) for all samples collected in October compared to in June, 208 including in soil samples, and estimated richness did not increase in the samples collected 209 between storms (Figure 6). Community compositions in field samples from the different months

210 segregated along the primary axis in PCA and the differences were statistically significant using

211 PERMANOVA (p < 0.01, F = 5.37, $R^2 = 0.17$), demonstrating that nearshore bacterial

communities were distinct in June and October (Supplemental Figure 7). Furthermore, soil
addition to October mesocosms did not cause observable increases in relative abundances of
phyla found in soil samples as it had in June (Figure 4, Supplemental Figure 8).

215 In order to identify individual taxa that became more abundant during storms and, thus, 216 contributed to the broader patterns observed in the data, we performed pairwise testing for 217 differentially abundant ASVs between sampling dates in each month. The greatest number of 218 significantly differentially abundant ASVs were found in the pairwise test between June 13 219 (before the storm) and June 16 (during the storm), with the vast majority being more abundant 220 during the storm (Figure 7A). In contrast, very few ASVs were significantly differentially 221 abundant in samples collected before and after the June storm (Figure 7A). While many of the 222 ASVs that had significantly higher relative abundance during the storm on June 16 were also 223 detected in soil samples, the majority were not (Figure 7A, B). ASVs that had significantly higher 224 relative abundance on June 16 compared to June 13 belonged to a total of 67 orders-including 225 Flavobacteriales, Campylobacterales, and Vibrionales, which can be pathogenic to humans and 226 marine organisms (Pruzzo et al., 2005; Silva et al., 2011; Loch and Faisal, 2015; Canty et al., 227 2020), and Rhizobiales, Sphingomonadales, and Alteromonadales, which were also found in 228 soil samples (Figure 7B). In pairwise tests for October samples, the largest number of 229 differentially abundant ASVs were likewise found between samples collected before (Sept. 28) 230 and between storms (Oct. 1 and 3), but none of the differentially abundant ASVs were also 231 found in soil samples (Figure 7C, D). ASVs showing significant differences in relative 232 abundance between samples taken before and between the October storms belonged to 21 233 different orders, with most, but not all, also represented in the June results (Figure 7D). 234

235 Discussion

237 As the severity and frequency of extreme storm events increases with global climate change, it 238 is increasingly important to understand how these events impact ecological functioning in 239 marine ecosystems (Wetz and Paerl, 2008; Du et al., 2019). However, characterizing the effects 240 of extreme storm events, such as typhoons, on coastal ecosystems is a complicated task, due 241 both to forecast unpredictability, which makes sampling before and after events challenging, 242 and the dangerous conditions that accompany storms (Chen et al., 2018). This study reports on 243 the nearshore microbial community dynamics and relevant environmental parameters during 244 two short sampling campaigns encompassing major storms of the 2018 typhoon season in 245 Okinawa, Japan. In addition, concurrent, controlled mesocosm experiments were performed to 246 supplement field observations and isolate impacts of terrestrial sediment input that regularly 247 accompanies large storms in Okinawa. Predictably, storms caused influxes of both freshwater 248 and sediment into the coastal marine environment (Figure 3, Supplemental Figure 1), which 249 carried some soil-derived bacteria and some bacteria presumably derived from other terrestrial 250 sources (e.g. urban or agricultural wastewater; Figures 4 and 7). The storm effects were short-251 lived, however, and bacterial community compositions in samples collected before and after 252 storms were largely similar (Figures 4, 5, 7). While field samples were collected three days after 253 storm events, mesocosm experiments showed that bacteria introduced with soil additions 254 diminished in relative abundance just four hours after soil input and were no longer detectable 255 24 hours later. Despite the rapidity of community shifts, terrestrially derived bacteria and marine 256 bacteria that increased in relative abundance during and after storms may still influence coastal 257 biogeochemical cycling and, in some cases, could be detrimental to the coastal ecosystem or 258 dangerous to human health. For instance, Rhodobacteraceae bacteria (order Rhodobacterales), 259 which became relatively more abundant during storms (Figure 7), are known to rapidly and 260 competitively exploit transient sources of particulate and dissolved organic matter and thus may 261 enhance remineralization in nearshore waters (Buchan et al., 2014). Other bacteria that 262 increased in relative abundance during and after storms can cause diseases in reef-building

263 corals (e.g. Vibrio spp., Pseudoalteromonas sp.) and humans (e.g. Campylobacter,

Fusobacteria, Enterobacter) (Pruzzo et al., 2005; Vizcaino et al., 2010; Silva et al., 2011;

Zimmer *et al.*, 2014); Figure 7, Supplemental Tables 6 and 7), emphasizing the need to better

266 understand sources and sinks of these bacteria in the coastal environment.

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Extreme storms cause rapid changes in environmental conditions and bacterial community composition

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271 During the storm event in June, we measured decreased salinity and increased turbidity in 272 nearshore surface waters (Figure 3), which is similar to previous reports following major storm 273 events (De Carlo et al., 2007; Zhou et al., 2012; Chen et al., 2018). Freshwater and sediment 274 input was accompanied by moderate increases in NO₂⁻, NO₃⁻, NH₄⁺, PO₄³⁻, SiO₂ and dFe 275 concentrations that ranged from a few nM, in the case of dFe, to up to 25 μ M for SiO₂ (Figure 3) 276 and is consistent with terrestrial run-off from agricultural areas with high iron-content soil, as is 277 found in Okinawa (Arakaki et al., 2005). However, nutrient loading dissipated before samples 278 were collected again three days after the June storm (Figure 3). Bacterial diversity (ASV 279 richness, Figure 6) increased during the storm, representing the introduction of bacterial taxa 280 that were also found in our soil samples and are common components of soil microbiomes, 281 including Acidobacteria, Actinobacteria, Chloroflexi, Firmicutes, Planctomycetes, Rokubacteria 282 and Verrucomicrobia (Figure 4; Freitas et al., 2012; Witt et al., 2012; Balmonte et al., 2016; Lin 283 et al., 2019). Many of these taxa are also commonly associated with particulate matter in marine 284 environments (e.g. Planctomycetes, Actinobacteria, and Verrucomicrobia) making resuspension 285 of marine sediment during storms another potential contributing source of these bacteria 286 (DeLong et al., 1993; Crespo et al., 2013). Campylobacterales (Epsilonbacteraeota) that 287 increased in relative abundance during storms could also derive from agricultural waste being incorporated into run-off (Jones, 2001). 288

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290 Three days after the storm, on June 19, the bacterial diversity returned to pre-storm levels 291 (Figure 6A) and the community composition was almost identical to before the storm (Figure 292 5A). Moreover, only 19 ASVs were significantly differentially abundant between June 13 and 293 June 19 compared to 267 ASVs between June 13 and June 16 (Figure 7A), demonstrating the 294 high speed at which the microbial community recovered. This high speed largely contrasts with 295 previous field studies where longer periods of time (i.e. from weeks to months) were needed to 296 restore pre-storm microbial compositions (De Carlo et al., 2007; Yeo et al., 2013). The 297 ephemerality of terrestrially-derived bacteria after introduction to the coastal environment was 298 similarly apparent in results from the June mesocosm experiment. We detected phyla dominant 299 in soil samples (e.g. Acidobacteria, Chloroflexi, Firmicutes, and Planctomycetes) in treatment 300 mesocosms one hour after soil was added, but the relative abundances of these phyla 301 decreased in samples taken four and twelve hours later, and they were no longer detectable in 302 samples taken 24 hours later. The transience of storm-effects on nearshore bacterial 303 communities was further demonstrated in results from the October storm event. The bacterial 304 community recovered just three days after super-typhoon Kong-Rey passed, the second super-305 typhoon to affect Okinawa in less than a week (Figure 5B). This transience may, in part, reflect 306 terrestrial and freshwater bacteria lysing in seawater due to osmotic stress. However, enteric 307 bacteria are able to survive in seawater, particularly when organic material is readily available 308 (Munro et al., 1989). Settling was likely also a contributing factor, as many of the bacterial taxa 309 that increased in relative abundance during storms and after soil addition to mesocosms are 310 known to associate with sediment surfaces (DeLong et al., 1993; Duret et al., 2019) and some 311 sedimentation was visible in mesocosm bottles despite using pumps to maintain water 312 circulation. Moreover, tidal flushing may have transported introduced or resuspended bacteria 313 offshore.

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315 Extreme storms cause context-dependent changes in environmental conditions and 316 bacterial community composition

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318 While storm effects were transient in both cases (Figure 5), the June and October storm 319 events affected coastal physicochemistry and bacterial community composition differently 320 (Figures 3, 4). Events in both June and October increased coastal nitrogen loading, but only the 321 June event was accompanied by increased PO₄³⁻ and SiO₂ concentrations, and the June storm 322 caused more pronounced shifts in bacterial community composition, highlighted by a larger 323 increase in ASV richness and a higher number of differentially abundant ASVs (Figure 7). 324 Although these differences in both nutrient loading and bacterial communities, may reflect 325 disparate sampling schemes, it is likely that differences in rain and wind intensity, as well as the 326 contexts in which the two events occurred, had strong effects. It has been shown that wind 327 speed alone can drive alpha diversity shifts in epipelagic bacterial communities (Bryant et al., 328 2016). Key differences between the June and October events were that much more rain, wind 329 and wave action accompanied the October storm event than the June storm and that the June 330 storm made landfall at the beginning of the typhoon season, whereas the October storms 331 affected Okinawa towards the end of the typhoon season.

332 Typhoon Trami, on Sept. 29, caused twice as much precipitation as tropical storm 333 Gaemi, on June 16 (Figure 2), which could have diluted nutrient loading in storm run-off and 334 caused the smaller changes in nutrient concentrations recorded in October compared to in June (Figure 3). When flushing rate is high, less PO_4^{3-} is desorbed from sediments and there is a 335 336 dilution effect for both dissolved phosphorus and nitrogen in run-off (Blanco et al., 2010). 337 Additionally, Gaemi was the first major storm of the 2018 typhoon season (June–October) 338 following a relatively prolonged dry period, which could increase nutrient loading in storm run-339 off, especially since agricultural fertilizers are applied throughout the preceding spring and 340 summer growing season. By October, several tropical storms and typhoons had already

341 affected Okinawa (Figure 2); the 2018 Pacific typhoon season had higher than average storm

342 frequency and included 29 tropical storms, 13 typhoons, and 7 super typhoons, although not all

343 made direct landfall with Okinawa (Japan Meteorological Agency,

344 <u>https://www.jma.go.jp/jma/indexe.html</u>). These intervening events could strip topsoil and deplete

soil nutrients and microbiomes, so that October storm run-off carried less nutrients, organic

346 material, and terrestrial bacteria into coastal water than storm run-off in June. Moreover,

347 antecedent soil moisture affects dissolved nutrient loading in run-off, with less nutrients

348 desorbing from clay-based soils, like Okinawa red soil, when already wet (Perrone and

349 Madramootoo, 1998).

350 The controlled mesocosm experiments offer additional insight for interpreting differences 351 in field observations between June and October. Despite collecting soil from the same place in 352 June and October, the soil microbiome in October had half the bacterial richness as in June 353 (Figure 6) and soil addition to October mesocosms did not introduce soil bacteria or increase 354 bacterial richness as it had in June mesocosms (Figure 4, Supplemental Figure 9). These 355 results suggest that run-off from storms occurring towards the end of the typhoon season may 356 carry less diverse bacterial assemblages than run-off from early-season storms. Furthermore, 357 soil addition in October mesocosms did not cause nitrogen (NO₂⁻ or NO₃⁻) or dFe to increase 358 over baseline measurements, but did cause increases in SiO₂ and PO₄³⁻ (Supplemental Figure 359 6). However, the increase in SiO₂ and PO_{4³⁻} were gradual, demonstrating that time is required to 360 release these compounds from red soil (De Carlo et al., 2007; Blanco et al., 2010; Chen et al., 361 2018). Soils in Okinawa could, therefore, have lower nitrogen content at the end of the typhoon 362 season and more intense rains may deliver less nutrients due to rapid flushing.

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364 Influence of extreme storms on bacterial taxa in nearshore waters

366 The influence of June and October storms on coastal microbial communities varied in effect 367 size, but in both instances the community shift was rapid and transient and included shared 368 bacterial groups that became more abundant during and after storms in both months: 15 out of 369 the 21 bacterial orders encompassing ASVs that were significantly more abundant during the 370 October event were shared with the June event (Figure 7). The shared bacterial orders included 371 primarily heterotrophic bacteria, several of which contain potentially pathogenic groups (e.g. 372 Vibrionales, Fusobacteriales, and Campylobacterales) (Pruzzo et al. 2005; Silva et al. 2011; 373 Canty et al. 2020). Originally, we expected cyanobacteria to respond to increased inorganic 374 nutrients delivered to the coastal ecosystem with storm run-off. Specifically, we expected 375 nitrogen fixers—such as Trichodesmium, which occasionally blooms near Okinawa (Grossmann et al., 2015)—to benefit from increased dFe and PO₄³⁻ in run-off (Sañudo-Wilhelmy et al., 2001; 376 377 Wu et al., 2003). Instead, only a few cyanobacteria ASVs became more abundant during the 378 June storm, but not the October storm (Figure 7). Namely, cyanobacteria belonging to the family 379 Oscillatoriaceae, which form filamentous benthic mats (Siegesmund et al., 2008; Engene et al., 380 2018), and Obscuribacterales, which are uncultured and have unknown morphology, but may 381 not be photosynthetic (Soo et al., 2014), became more abundant in June. 382 There are several possible explanations for the minimal changes in cyanobacteria relative 383 abundance: i) cyanobacteria were not limited by compounds present in run-off, ii) nutrients in 384 run-off were not biologically available, or iii) terrestrial run-off was transported offshore or diluted 385 too guickly to affect nearshore communities. Nutrient concentrations that were elevated during 386 the June storm (NO₂, NO₃, PO₄³ and dFe) decreased after the storm passed, but not 387 immediately and not as quickly as microbial communities recovered (Figures 3, 4). Since these 388 nutrients were not immediately drawn down, the photosynthetic microbial community may not 389 have been nutrient limited in our study area. That microbial communities did not respond to soil addition in mesocosms, despite increased PO₄³ and dFe, further suggests that nutrient 390 availability was not driving major changes in bacterial community composition. 391

392 Heterotrophic bacteria increasing in relative abundance during and after storms could derive 393 from soil or other components of run-off, have been resuspended from the seafloor due to wind 394 and waves, or have responded to increased organic matter from run-off. These bacteria may 395 affect ecosystem functioning by remineralizing organic matter within the near-shore environment 396 (Buchan et al. 2014; Kieft et al. 2018) or by causing opportunistic infections in keystone 397 metazoans (Pruzzo et al. 2005; Silva et al. 2011; Canty et al. 2020), thus making it important to 398 identify their source. Rare bacteria often become more abundant during or after disturbances 399 (Preisner et al. 2016; Fuentes et al. 2016; Sjöstedt et al. 2012), but it is difficult to parse out their 400 sources (Shade et al. 2014). Such conditionally rare bacteria may result from introductions or 401 opportunistic growth of previously undetectable bacteria (Shade et al. 2014). Discriminating 402 between these potential sources is challenging with current methods, but prolonged intensive 403 sampling better resolves the dynamics of rare taxa (Shade et al. 2013; Shade et al. 2014). 404 Beyond working towards establishing the more long term dynamics of nearshore 405 bacterioplankton communities, future work aimed at characterizing bacteria in wastewater 406 outflows and stormwater could also help determine whether bacteria that become more 407 abundant during storms represent introductions or expansions. 408 409 Potential ecosystem consequences due to terrestrial run-off from extreme storms 410 411 Despite being short-lived, the changes we observed in bacterial community composition and

environmental parameters during storms can nevertheless be detrimental to both the coastal
ecosystem and human health. While most terrestrially-derived bacteria are benign to marine
ecosystems, many are potentially pathogenic to corals and other marine organisms (Sutherland *et al.*, 2004; Haapkylä *et al.*, 2011; Pollock *et al.*, 2014; Sheridan *et al.*, 2014). Terrestrial run-off
has been implicated in coral diseases, such as White Pox Disease and Red Band Disease, in
many tropical and subtropical locations, including Madagascar (Haapkylä *et al.*, 2011; Pollock *et al.*

418 al., 2014; Sheridan et al., 2014), the Caribbean (Frias-Lopez et al., 2002; Patterson et al., 419 2002), and Australia's Great Barrier Reef (Pollock et al., 2014). Furthermore, storm events can 420 cause water-column mixing and sediment resuspension, leading to pathogenic marine bacteria 421 that usually inhabit the seafloor to become more abundant in the water column (Hassard et al., 422 2016). Indeed, in our study we found several strains of Vibrio spp., Pseudoalteromonas sp., and 423 Rhodobacteraceae bacteria specifically associated with coral disease (Sussman et al., 2008; 424 Sheridan et al., 2014) to be significantly enriched in samples we collected during or following 425 storm events (Supplemental Table 5). Considering the additional stress caused by turbidity and 426 sedimentation during storm events, alongside potentially decreased pH and dissolved oxygen 427 due to enhanced heterotrophic bacterial respiration (Weber et al., 2012; Altieri et al., 2017), 428 corals and other organisms may be especially susceptible to pathogen infection during and after 429 extreme storms.

430

431 Heavy rains and floods have long been implicated in transporting human pathogens (e.g. fecal 432 coliform bacteria) to the marine environment (Pandey et al., 2014; De Jesus Crespo et al., 433 2019). Indeed, we found bacterial taxa that are potentially dangerous to humans-including 434 Campylobacterales, Fusobacteriales, Bacillales, Clostridiales and Enterobacteriales (Bennett 435 and Eley, 1993; Sharma et al., 2003; Davin-Regli et al., 2019)-significantly enriched during and 436 following storms. Many of these bacteria were not detected in soil samples, suggesting 437 additional sources of bacterial contamination in storm run-off. For instance, some of the 438 Bacillales and Enterobacterales ASVs enriched in storm-influenced water samples were not 439 present in soil samples and none of the enriched Clostridiales, Campylobacterales and 440 Fusobacteriales ASVs were also found in soil samples (Figure 7). Likewise, these taxa were not 441 found in mesocosms following soil addition in June or October, demonstrating the larger effect 442 of storms and run-off on the coastal ecosystem than simply transporting sediment into the water. Human pathogens may derive from live-stock, storm drains, or overwhelmed waste 443

444 treatment plants during storms (Weiskel et al., 1996; Jamieson et al., 2004). Interestingly, these 445 taxa were already present in water samples collected on September 28-although they were 446 more abundant on October 1 and 3-further indicating a cumulative effect of the typhoon 447 season on the coastal ecosystem and emphasizing the context-dependency of storm effects. 448 Ultimately, swimmers and other recreational users need to be aware that pathogenic bacteria 449 are likely present in Okinawa coastal waters following large rain events. This is not currently the 450 case in Okinawa as it is in other locations with more robust coastal monitoring programs in 451 place (e.g. the Southern California Water Research Project or the DNAgua-Net European 452 project).

453

454 **Conclusions & Future Directions**

455

456 Despite challenges associated with sampling marine ecosystems during tropical storms and 457 typhoons, this study describes the timing and nature of storm effects on coastal bacterial 458 communities in Okinawa, Japan. We found that storm effects were transient, but highly context-459 dependent. We coupled controlled mesocosm studies with field observations in an effort to 460 disentangle the effects of extreme wind and waves and enhanced currents from the effects of 461 soil input during storms. While the mesocosm results were useful, future studies would benefit 462 from more realistic run-off simulation (e.g. Le et al., 2016) than the soil additions we employed. 463 Furthermore, it remains that we did not perform bacterial cell counts or otherwise measure 464 bacterial biomass or metabolic activity, thus leaving the possibility that run-off increased 465 microbial biomass or differentially affected microbial physiology. Future studies may capture 466 such responses by measuring bacterial respiration rates or by performing metatranscriptomics. 467

468 It is important to note that environmental effects of extreme storms will vary in terms of intensity,
469 spatial extent and duration in different ecosystems and need to be evaluated locally (Paerl *et al.*,

2006; Zhang *et al.*, 2009; Herbeck *et al.*, 2011). Storm effects were transient in the open, tidallyflushed Okinawa coast, but more prolonged storm effects have been observed in other coastal
systems, particularly semi-enclosed areas, such as bays and estuaries, where terrestrial
sediment loads can have residence times from weeks to years (Paerl *et al.*, 2001; Zhang *et al.*,
2009; Herbeck *et al.*, 2011). Therefore, we suggest that the short-term study of typhoon events
follow an adaptive sampling strategy (Wetz and Paerl, 2008), which involves the definition of
well-established baselines for various physicochemical and biological parameters.

477

478 Finally, the transient nature of storm effects described here should not be viewed as diminishing 479 their potential impact on reef or human health. Differentially abundant bacteria during storms 480 may cause disease in marine organisms and humans. With future climate change scenarios 481 predicting more frequent and destructive storms and continued expansion of tourism and 482 agriculture activities, it is likely that the amount of terrestrial run-off and associated risks will 483 escalate in the future. This makes regional monitoring programs, including a comprehensive 484 understanding of background conditions, essential for better interpretations of ecological 485 consequences from extreme storm events.

486

487

- 488 **Experimental Procedures**
- 489

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490 Study setting
```

491 Seawater was collected for metabarcoding and physicochemical analysis from four nearshore

492 sampling points approximately 250–500 m apart, along the central west coast of Okinawa

493 Island—a semi-urban region with mixed land-use, including agriculture and coastal development

- 494 projects (Figure 1). The sampling points (A1–4) were each at least 1.2 km from the nearest
- 495 concentrated fresh-water input (e.g. streams or rivers). At the start of the 2018 typhoon season,

samples were collected before (June 13), during (June 16), and after (June 19) tropical storm
Gaemi, which struck Okinawa on June 16, 2018 and caused a red soil pollution event (Figure 2,
Supplemental Figure 2A). Towards the end of the typhoon season, samples were collected
before (Sept 28), during (Oct 1 and Oct 3) and after (Oct 8) a red soil pollution event caused by
typhoon Trami, which made landfall with Okinawa on September 30, and was prolonged by
Typhoon Kong-Rey, which approached Okinawa on October 5th (Figure 2, Supplemental Figure
2B–C).

503

504 Seawater sampling for DNA and physicochemical analysis

505 Surface seawater was collected for DNA metabarcoding by submerging clean 500 mL Nalgene 506 bottles just below the sea surface. Seawater for dissolved Fe (dFe) and nutrient analysis was 507 collected in acid-cleaned 50 mL Falcon tubes. Physicochemical properties-dissolved oxygen 508 (DO), salinity, temperature, and turbidity—were measured with a CTD probe (RINKO, JFE 509 Advantech, Japan) at each site. After being immediately transported to the laboratory, seawater 510 samples for metabarcoding were filtered through 0.2 µm pore-size Polytetrafluoroethylene 511 (PTFE) filters (Millipore) under gentle vacuum and filters were stored at -20 °C for later DNA 512 extraction. Seawater samples for dFe and nutrient analysis were filtered through 0.45 µm pore-513 size acid-washed Teflon digiFILTERS (SPC Science, Canada) and filtered water samples were 514 stored at -20 °C for later chemical analysis.

515

516 Mesocosm experimental design

517 Seawater was collected for concurrent mesocosm experiments on June 11 (26.512 °N,

518 127.872°E) and October 10 (26.479 °N, 127.829 °E). Nearshore coastal seawater was pumped

519 from just below the sea surface and filtered through 1 mm and 300 µm nylon mesh screen sizes

- 520 to remove debris and larger organisms. Acid-cleaned 22 L clear-plastic carboys were rinsed
- 521 twice with 300 μm nylon mesh-size filtered seawater before filling to 20 L with filtered seawater.

522 Bottles were covered with parafilm and kept shaded during transport to the Okinawa Institute of Science and Technology (OIST) Marine Science Station, where they were submerged in a basin 523 524 with continuous flow-through seawater to keep conditions within the bottles similar to natural 525 conditions. Mesocosm bottles were topped with silicone sponge stoppers to allow gas 526 exchange, but limit evaporation and prevent dust, water or other contaminants from entering the 527 bottles during the experiment (Supplemental Figure 3). Small pumps were included in each 528 mesocosm to maintain water circulation (2 L min⁻¹). HOBO temperature and light loggers 529 (Onset) were fastened to the pumps and at the same depth in the basin surrounding 530 mesocosms to ensure that mesocosms conditions remained similar to ambient conditions 531 (Supplemental Figure 4). Salinity and DO were measured each time water was sampled from 532 the mesocosms throughout the experiment (Supplemental Figure 5). Mesocosms experiments 533 included 9-10 mesocosms: 4-5 control replicates (four in June and five in October) and 5 534 treatments replicates with red soil added to an ecologically relevant high concentration of 200 535 mg L⁻¹ (O'Connor et al., 2016). Mesocosms were sampled (100 mL for metabarcoding, 50 mL 536 for dFe and nutrient concentration) with 50 mL sterile pipettes before the experiment started 537 (t0), and 1, 4, 12, 24, and 48 h following red soil addition to treatment bottles. Water samples 538 were processed as described in the previous section.

539

540 **Red soil collection for mesocosm experiments and DNA analysis**

Soil samples were collected from an open agricultural field with exposed soil (26.507 °N, 127.868 °E), on June 10 and October 9 for addition to red soil treated mesocosms and to evaluate soil microbiomes. Soil samples were sieved through 330 µm mesh and maintained at 4 °C for 24 h until use in the mesocosm experiment. To determine soil moisture content, 10 g subsamples (n=10) were weighed and dried at 100 °C by following the standard method AS 1289.2.1.1-2005 (Standards Association of Australia). Soil moisture content was used to calculate how much wet soil should be added to mesocosms in order to reach the final

concentration of 200 g soil (dry weight) per L seawater. In both June (n=4) and October (n=2),
additional 50 g aliquots of soil were kept at -20 °C for subsequent metabarcoding analysis.

550

551 Chemical analysis: dFe and major nutrients

552 Dissolved Fe (dFe) concentration was determined following the methodology of Wu and Boyle 553 (1998). This method uses a Mg (OH)₂ co-precipitation to pre-concentrate Fe from seawater 554 followed by an isotope dilution method. Seawater dFe was guantified using an internal standard 555 element (⁵⁷Fe) with inductively coupled plasma mass spectrometry (Element 2, Thermo 556 Scientific). The mass spectrometer was operated in medium-resolution mode with 4000 557 resolution (FWHM). The mass calibration was performed using a multi-element ICP-MS tune-up 558 solution (Thermo Fisher Scientific). In order to ensure the quality of the ICP-MS analysis, control 559 standards and samples (i.e. analytical replicates, certified reference material and analytical 560 blank) were analyzed once every 12 samples. Recovery of Fe from reference certified material 561 QC3163 (Sigma-Aldrich, USA) was satisfactory and ranged from 65 to 80%. The overall error 562 associated with the analytical process was typically lower than 5% and never higher than 15%. 563 All measurements were above the instrument's detection limit. Analysis was carried out at the 564 OIST Instrumental Analysis Section mass spectrometry laboratory. Special attention was paid to 565 avoid Fe contamination and an exhaustive cleaning process was carried out following the 566 methods of (King and Barbeau, 2011).

567

Nutrient concentrations—including Nitrate (NO₃⁻), Nitrite (NO₂⁻), Ammonium (NH₄⁺), Phosphate
(PO₄³⁻) and Silica (SiO₂)—were determined on a QuAAtro39 Continuous Segmented Flow
Analyzer (SEAL Analytical) following manufacturer guidelines. Final concentrations were
calculated through AACE software (SEAL Analytical). Nutrient Analysis was carried out at the
Okinawa Prefecture Fisheries and Ocean Technology Center.

574 Nutrient and dFe statistical analyses

575 Mesocosm data were found to be normally distributed with a Shapiro-Wilks test and, therefore, 576 a one-way ANOVA was performed to test overall differences between treatments. Post-hoc 577 Tukey HSD analysis was performed to identify which specific groups differed. Field data were 578 not normally distributed, regardless of transformation, so a Kruskal-Wallis test was used to test 579 for significant differences between sampling dates. Analyses were performed within the R 580 statistical environment (R Core Team 2018).

581

582 DNA extraction and metabarcode sequencing

583 DNA was extracted from frozen PTFE filters following the manufacturer protocol for the DNeasy

584 PowerWater Kit (Qiagen), including the optional heating step. DNA was extracted from soil

samples by following manufacturer protocol for the DNeasy PowerSoil Kit (Qiagen).

586 Metabarcode sequencing libraries were prepared for the V3/V4 region of the bacterial 16S

ribosomal RNA gene following Illumina's "16S Metagenomic Sequencing Library Preparation"

588 manual without any modifications. Sequencing libraries were transferred to the OIST

589 Sequencing Center for 2x300-bp sequencing on the Illumina MiSeq platform with v3 chemistry.

590 Overall, 18.4 million sequencing reads were generated in this study, with 76,217–219,584

sequencing reads per sample (mean = 137,585). Sequencing data are available from the NCBI

592 Sequencing Read Archive under the accession PRJNA564579.

593

594 Metabarcode analyses

595 Sequencing reads were denoised using the Divisive Amplicon Denoising Algorithm (Callahan et

596 *al.*, 2016) with the DADA2 plug-in for QIIME 2 (Bolyen *et al.*, 2019). Following denoising, 11.8

597 million sequences remained with 11,061–153,646 sequences per sample (mean = 88,033). A

total of 36,007 ASVs were identified in our dataset, with 64–2,886 unique ASVs observed per

599 sample (mean = 642). Taxonomy was assigned to representative ASVs using a Naive Bayes 600 classifier trained on the SILVA 97% consensus taxonomy (version 132, Quast et al., 2013) with 601 the QIIME 2 feature-classifier plug-in (Bokulich et al., 2018). The results were imported into the 602 R statistical environment (R Core Team 2018) for further analysis with the Bioconductor 603 phyloseq package (McMurdie and Holmes, 2013). The ASV richness for each sample was 604 estimated using the R package breakaway (Willis and Bunge, 2015) and differences in 605 estimated richness between sample types were tested for with the betta function (Willis et al., 606 2017). In order to minimize compositional bias inherent in metabarcoding data as much as 607 possible, we used the Aitchison distance between samples, which includes a centered log ratio 608 transformation to normalize data (Gloor et al., 2017), for principal component analyses (PCA). 609 Permutational analyses of variance (PERMANOVA) on Aitchison distances were performed with 610 the adonis function (999 permutations) in the R package vegan (Oksanen et al., 2019) to test 611 whether shifts in community composition were statistically significant. Lastly, we used the 612 DESeq2 bioconductor package (Love et al., 2014) to determine which ASVs were significantly 613 differentially abundant (False Discovery Rate adjusted *p*-value < 0.05) in water samples 614 collected from field sites before, during, and after storms. We then checked if significantly 615 differentially abundant ASVs were also present in soil samples from respective months to 616 assess whether differentially abundant ASVs were soil-derived. Intermediate data files and the 617 code necessary to replicate analyses are available in a GitHub repository:

- 618 <u>https://github.com/maggimars/RedSoil</u>.
- 619

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- 630
- 631

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- 904
- 905
- 906 Figure legends
- 907 Figure 1. Location of the study area A) Map locating Okinawa Island in the East China
- 908 Sea. B) Map showing the location of the central-west coast of Okinawa Island. C)
- 909 Location of the 4 nearshore sampling points in Onna-son, on the central-west
- 910 coast of Okinawa Island.
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Figure 2. Precipitation (in mm day⁻¹) and wind speed (m s⁻¹) during the 2018 912 913 typhoon season in Okinawa, Japan. Data were collected with the 914 meteorological station located at OIST Marine Science Station (26.510046 °N. 915 127.871721 °E) from June through October, corresponding to the duration of the 916 typhoon season in Okinawa. Bars represent the daily amount of rain (in mm); 917 dashed black and red lines indicate daily mean and maximum wind speeds (in m 918 s⁻¹); shaded areas represent the two red soil pollution events monitored in this 919 study: Tropical Storm Gaemi on June 16 and the super-typhoons Trami and 920 Kong-Rey on September 29 and October 5. The dates when water samples were 921 collected for chemical and DNA analyses are noted on the x-axis. 922 923 Figure 3. Temporal variation of physicochemical parameters before, during, and 924 after storm events in June and October, 2018. Bars represent the values of 925 temperature (°C), turbidity (NTU), salinity (‰) and concentrations of micro- and 926 macro-nutrient concentrations (NO₂⁻, NO₃⁻, NH₄⁺, PO₄³⁻, SiO₂ in µM; dFe in nM) at 927 sampling sites A1–4 along the central west coast of Okinawa, Japan. Error bars 928 represent one standard deviation of the mean from four replicates. Red dashed 929 vertical lines represent the timing of major storms and associated red soil 930 pollution events. Sampling dates when samples were also processed for 931 metabarcode analyses are indicated in bold on the x-axis. Temperature (°C), 932 turbidity (NTU) and salinity (‰) measurements were taken with a CTD probe. 933 Macro-nutrient concentrations were determined on a QuAAtro39 Continuous 934 Segmented Flow Analyzer and dFe concentration was determined by ICP-MS 935 after $Mg(OH)_2$ co-precipitation using the isotope dilution method(Wu and Boyle, 936 1998). 937

938 Figure 4. Relative abundance of bacterial phyla in field and mesocosm samples 939 collected in June and October, 2018. Each stacked bar represents the relative 940 contribution of major bacterial phyla to the total community at one sampling 941 location, in one red soil sample, or at one time point in mesocosms (mesocosm 942 bars represent aggregate data from 4–5 replicates depending on month and 943 treatment). Sampling stations (A)1-4 were sampled before, during, and after 944 major storms affecting Okinawa in June and October 2018. In June, the storm 945 affected Okinawa on 6/16, so that sampling on 6/13 was before the storm and 946 sampling on 6/19 was afterward. In October, two super typhoons affected 947 Okinawa on 9/28 and 10/05 so that 9/28 was before the event window, 10/01 948 and 10/03 was during the event window, and 10/08 was afterward. Red soil 949 samples are labeled with replicate numbers and were collected within a week of 950 the storms in June and October. Red soil (200 mg L⁻¹) was added to treatment 951 mesocosms immediately following the t0 sampling.

952

Figure 5. Principal component analysis (PCA) of Aitchison distances between
bacterial community composition before, during, and after storm events in
June (A) and October (B), 2018. (A) June 13 was before the June storm, 6/16
was during, and 6/19 was after. For the October storms (B), 9/28 was before the
storms, 10/01 and 10/03 were between, and 10/08 was after. Samples collected
during/between storm events separate from samples collected before and after
storms on PC1 for both the June and October events.

960

Figure 6. Richness estimates for bacterial communities in red soil samples and
 surface water samples collected before, during, and after storm events in
 June (A) and October (B), 2018. Red dashed vertical lines represent the timing

964	of major storms. In June (A), the red soil samples had significantly higher
965	richness than the water samples collected before and after the storm, but not
966	samples collected during the storm. In October (B), red soil samples had
967	significantly higher richness than all water samples and richness was not
968	significantly different in water samples collected before, during, or after the
969	storms. Differences in richness were considered significant when $p < 0.05$.
970	
971	Figure 7. Number of Amplicon Sequence Variants (ASVs) significantly
972	differentially abundant in pairwise comparisons between sampling dates in
973	June (A) and October (C) and log2 Fold Change of individual ASVs in
974	pairwise comparisons from before to during storms and from before to
975	after storms in June (B) and October (D), 2018. ASVs were considered
976	significantly differentially abundant when the FDR adjusted p -value was less than
977	0.05. ASVs that were also present in soil samples are colored red in panels A
978	and B; no differentially abundant ASVs in pairwise tests for October samples
979	were found to be present in October soil samples. Positive log2FoldChange
980	values (x-axis, B, D) indicate higher abundance of ASVs during/between storms
981	compared to before (J13 to J16, S28 to O01 + O03) or after storms compared to
982	before (J13 to J19, S28 to O08). Differentially abundant ASVs are grouped by
983	taxonomic order and orders are color coded by phylum so that colors correspond
984	to Figure 4.
985	
986	



Longitude







A. June







